

# Processing of Wave Directional Spectra into a climatology of swell events

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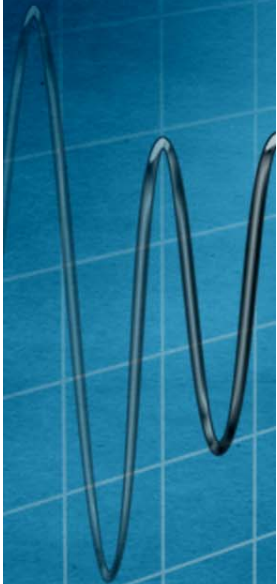
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« Time-series analysis in Marine science and applications for industry »

Conference in Logonna-Daoulas, France, 17-22 sept. 2012

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## Processing of Wave Directional Spectra



The data we  
have

What industry can  
use for design and  
operations

... into a climatology of swell events



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**Some applications require detailed knowledge of the long-term spectral wave climate at a given location.**



**Structural fatigue**



**Coastal erosion**



**Wave energy extraction**

**Etc...**

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**Problem: predict likely wave action over a long future duration, typically a few decades, when responses are sensitive to height, period and direction.**

If sea states can be represented by a single triplet  $H, T, \theta$ , and occurrence probability, and if their sequencing has neglectible influence:

Binning of the database parameters, selection of a few representative cases

Computation of action for those cases

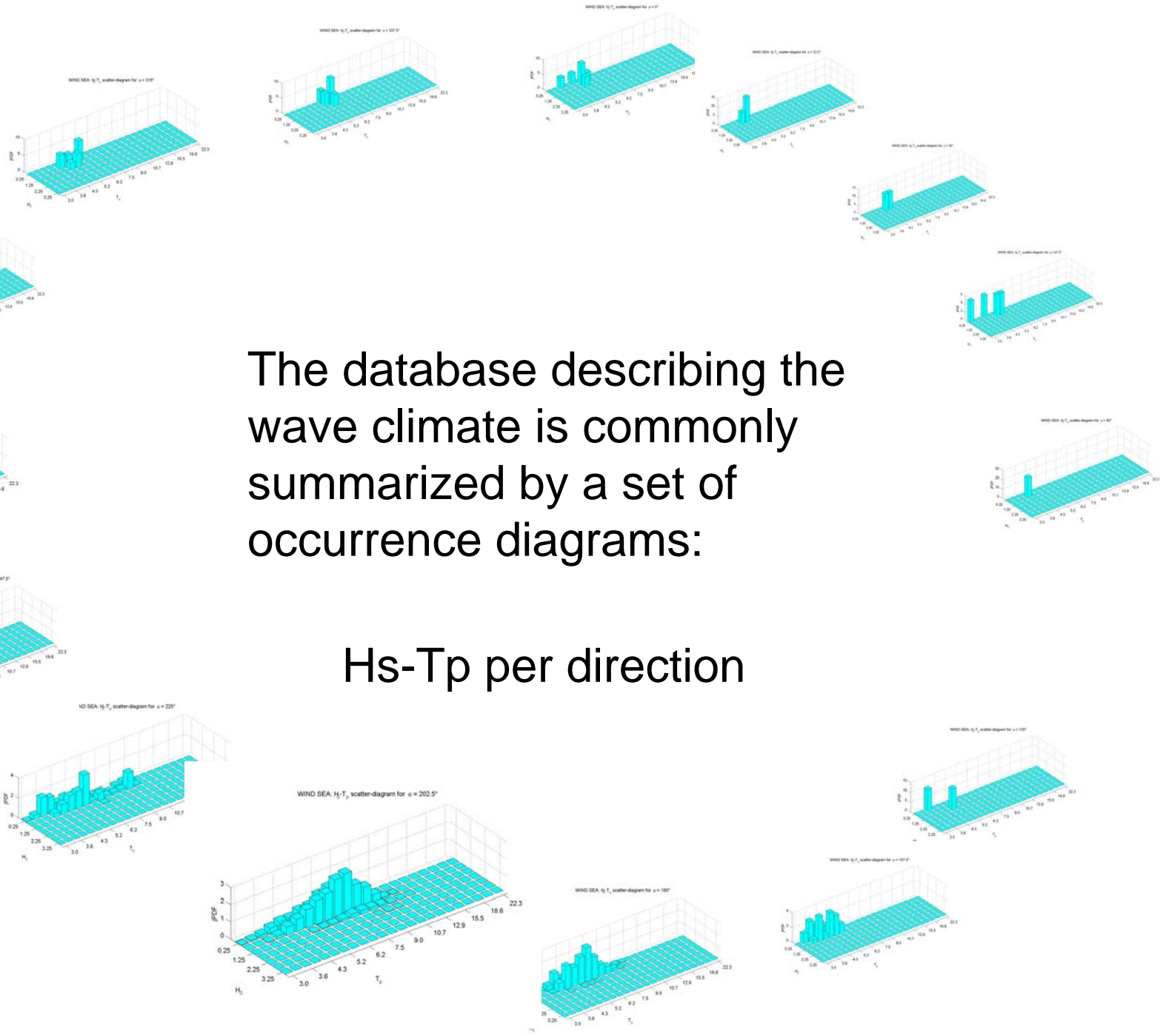
Estimation of the short term effects

Summation over the sea states, weighted by the occurrence probabilities

Order of magnitude:  
 $10^2$  to  $10^3$

For Fatigue: Simulations,  
RAOs, QTFs,  
FE models

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The database describing the  
wave climate is commonly  
summarized by a set of  
occurrence diagrams:

Hs-Tp per direction

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**On the other hand, directional measurements are costly to set up, and they often cover only short durations before they come to an end,**

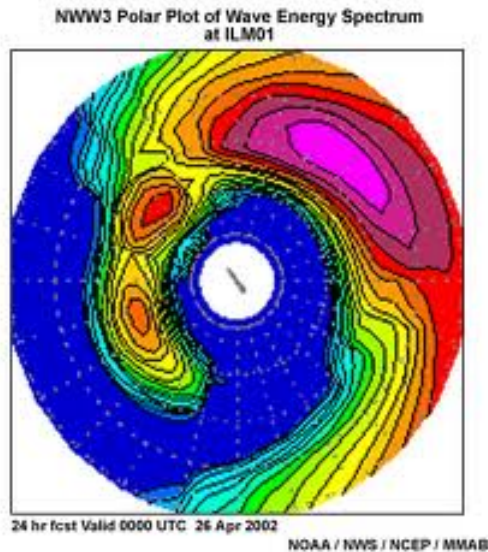


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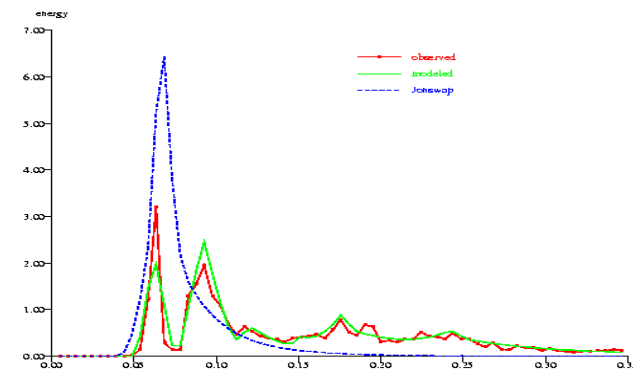
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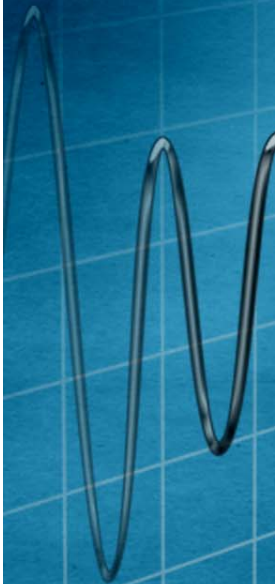


**...and at some locations, individual directional spectra can already not be characterized without a large number of parameters, let not say what it is for time-series of them !**



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If sea states can be represented by a single triplet  $H, T, \theta$ , and occurrence probability, and if their sequencing has neglectible influence:

Not enough data  
to estimate  
properly

Predictive value of  
study falls  
dramatically !

Commonly 3  
triplets, sometimes  
even more

Order of magnitude  
of the number of  
cases to consider  
raises to  $10^9$  !



# IO Illustrate Joint Probabilities estimation difficulty

8040 measured sea states

**SPOP** (existing partitionning tool)

8038 Main Swells  
5464 Secondary Swells  
4169 Wind Seas

Metocean  
specifications of the  
operator

# Naive reconstruction

	after partition		after assemblage	
	Number	Frequency [%]	Number	Frequency [%]
main swell	8038	99.97	8037	99.96
secondary swell	5464	67.96	5289	65.78
wind sea	4169	51.85	4088	50.85

Table 3: Proportions of the components.

	after partition		after assemblage	
	Number	Frequency [%]	Number	Frequency [%]
MS only	1158	14.40	1212	15.07
MS + SS	2713	33.74	2740	34.08
MS + WS	1416	17.61	1536	19.10
MS + SS+ WS	2751	34.22	2549	31.70
no MS (i.e. WS only)	2	0.03	3	0.05
Total	8040	100	8040	100

Table 4: Proportions of the combinations.

# ...does not provide the right sea states !

based on empirical statistics of the metocean database

	MS only	MS+SS	MS+WS	MS+SS+WS
$\overline{H_S}$ [m]	1.59	1.40	1.25	1.22
$\sigma_{H_S}$ [m]	0.48	0.40	0.32	0.32
$\chi_{H_S}$ [m <sup>3</sup> ]	0.95	1.01	1.16	1.19
$\kappa_{H_S}$ [m <sup>4</sup> ]	3.65	4.35	5.81	5.15
$H_{S,1/3}$ [m]	2.14	1.85	1.60	1.58
$H_{S,1/10}$ [m]	2.63	2.23	1.90	1.91
$T_p$ [s]	12.23	11.21	9.06	9.36
$\sigma_{T_p}$ [s]	2.00	1.51	1.63	1.82
$\chi_{T_p}$ [s <sup>3</sup> ]	0.15	0.45	0.23	0.14
$\kappa_{T_p}$ [s <sup>4</sup> ]	2.39	3.03	2.71	2.61
$T_{p,1/3}$ [s]	14.48	12.90	10.90	11.38
$T_{p,1/10}$ [s]	15.73	14.15	12.06	12.62
$\theta$	208.26	202.08	210.50	206.09
$\sigma_\theta$	8.84	7.05	10.47	8.91
$\chi_\theta$	0.38	0.24	1.16	0.55
$\kappa_\theta$	2.79	4.05	8.89	3.32
$\theta_{1/3}$	218.10	209.66	222.00	216.20
$\theta_{1/10}$	225.23	215.39	229.71	223.30

based on the new statistics from metocean specifications

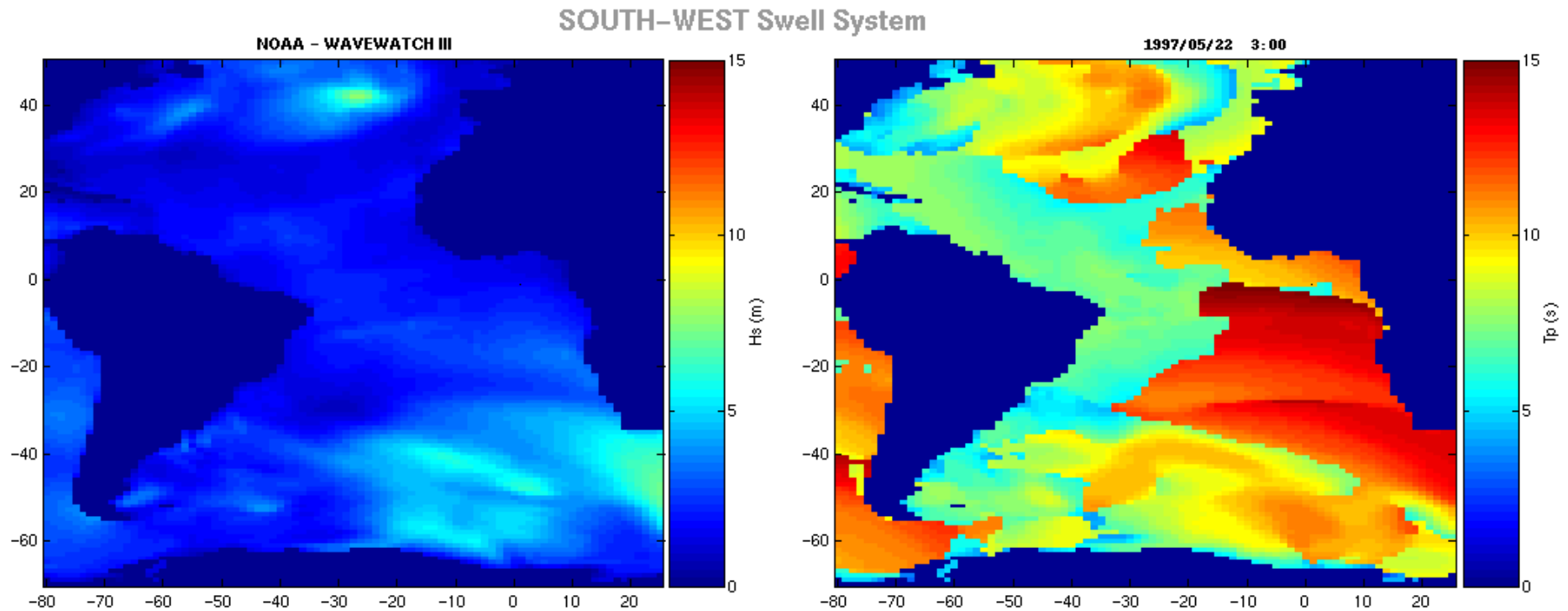
	MS only	MS+SS	MS+WS	MS+SS+WS
$\overline{H_S}$ [m]	0.99	1.27	1.20	1.43
$\sigma_{H_S}$ [m]	0.49	0.50	0.46	0.47
$\chi_{H_S}$ [m <sup>3</sup> ]	0.97	0.71	0.96	0.74
$\kappa_{H_S}$ [m <sup>4</sup> ]	4.65	3.96	4.88	4.13
$H_{S,1/3}$ [m]	1.25	1.46	1.27	1.64
$H_{S,1/10}$ [m]	1.75	1.90	1.77	2.05
$T_p$ [s]	11.08	11.72	8.69	9.45
$\sigma_{T_p}$ [s]	2.49	1.94	2.21	2.06
$\chi_{T_p}$ [s <sup>3</sup> ]	0.68	0.33	0.52	0.26
$\kappa_{T_p}$ [s <sup>4</sup> ]	3.40	3.03	3.06	2.80
$T_{p,1/3}$ [s]	11.78	12.49	9.49	10.28
$T_{p,1/10}$ [s]	14.13	14.30	11.70	12.21
$\theta$	205.80	205.96	205.76	205.92
$\sigma_\theta$	9.14	9.10	9.07	9.03
$\chi_\theta$	1.62	1.64	1.62	1.63
$\kappa_\theta$	7.93	7.97	7.90	7.95
$\theta_{1/3}$	202.5	202.82	202.5	202.82
$\theta_{1/10}$	225	225	224.73	224.82

Table 7: Statistical properties of combinations parameter.

Especially, single swells lead to highest  $H_s$  in measurements, to lowest in naive reconstruction.



# What makes the observed spectra's swell part ?



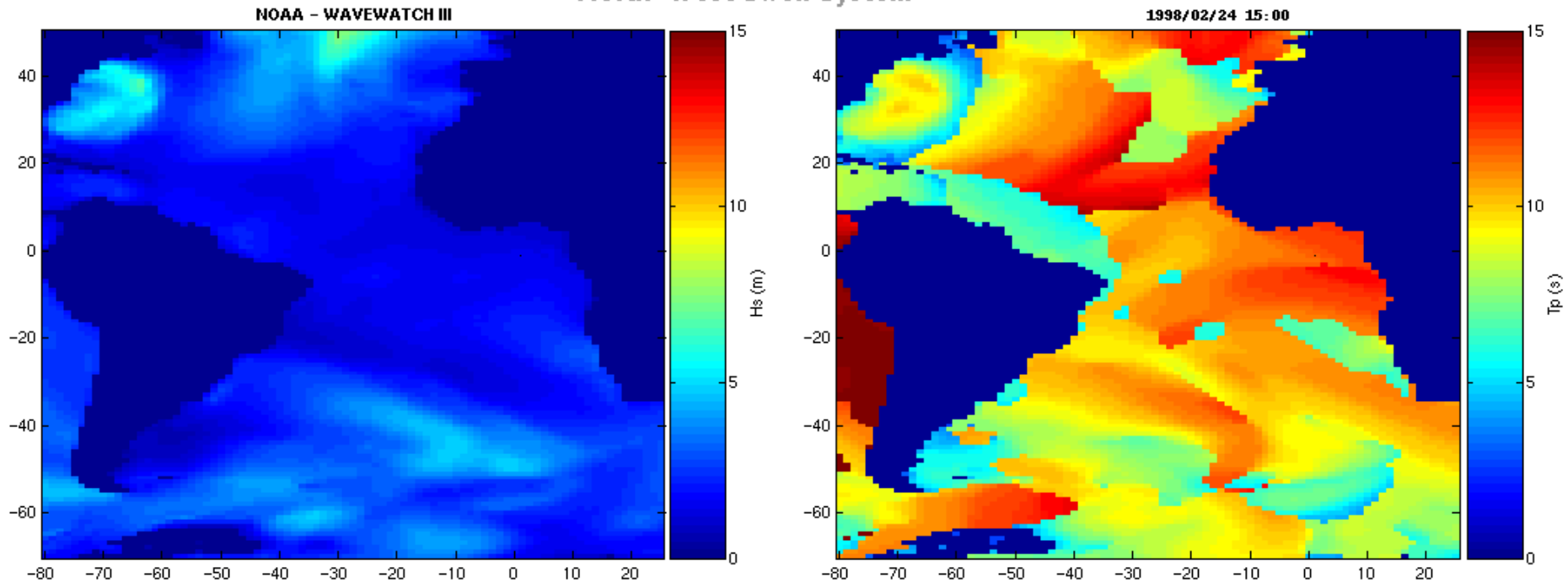
Let's consider swell arriving on West Africa oil fields (or endangered shorelines).

Left displays significant wave height, thus storms passing by in the roaring forties, and right dominant wave periods, thus the front of the overtaking of the existing swells by a new primary one sent by the storm.

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### North-West Swell System



In some conditions, swells from the North Atlantic may hit beamside structures heading to the southwesterly dominant waves: there is no way to construct some equivalent spectrum with single  $H_s$ ,  $T_p$ , direction for several swells present in a sea state.

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# Characteristics of W.A. spectra

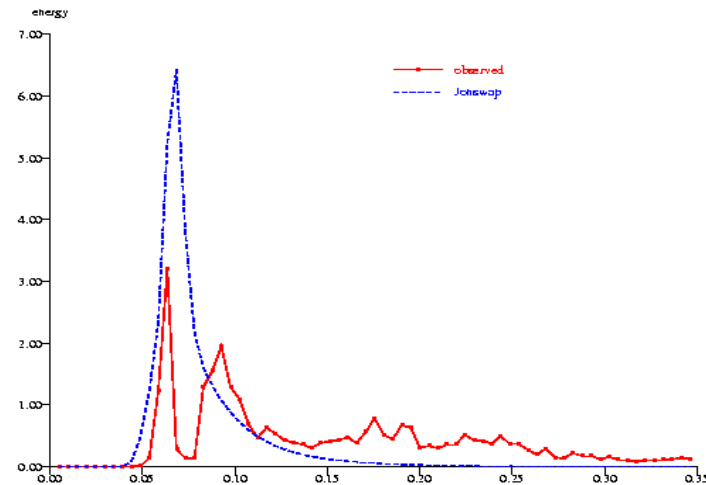
- Multiple swell peaks
- Deep troughs in between
- Still some wind sea





# Limitations of standard models

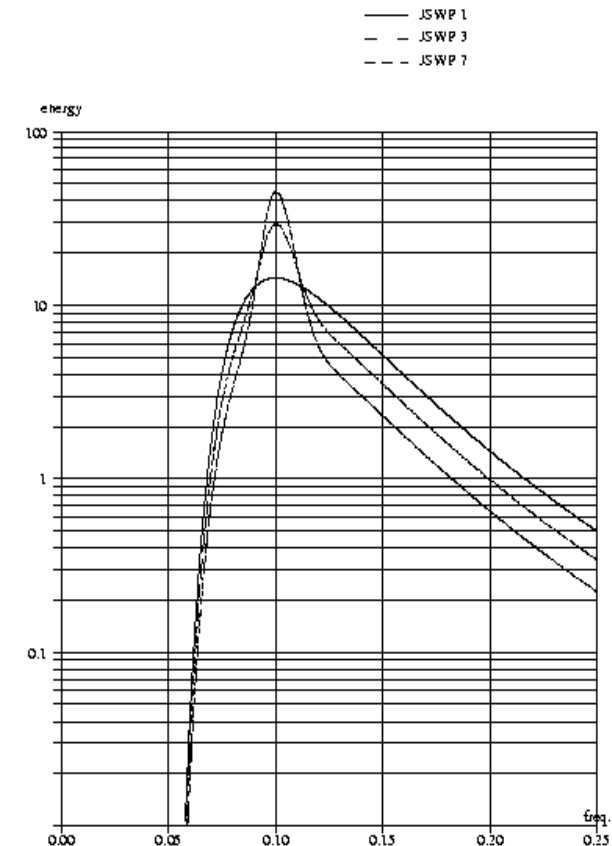
- 2 peaks at most.
- Spectral shapes for individual systems are fully or not fully developed WIND seas.
- Gaps between peaks poorly represented.
- High  $\gamma$  values, no physical meaning, numerical and sampling problems.



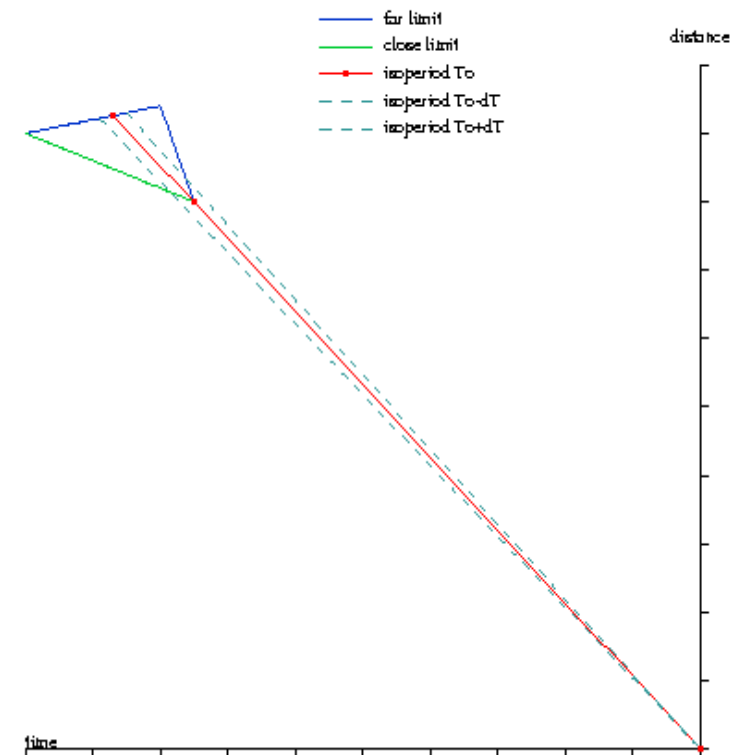
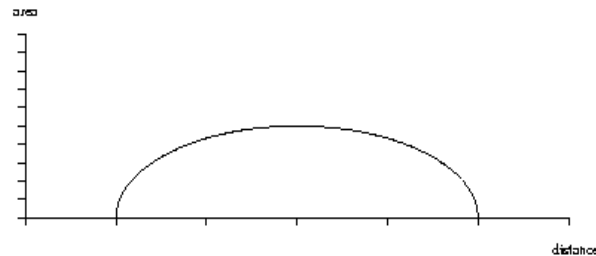
# Is JONSWAP suited to swell ?

**JONSWAP model means  
that wind has not blown  
enough to fill-in the  
missing part with respect  
to a P-M.**

**It is an enhancement of the  
model for wind to waves  
spectral energy transfer  
to account for what  
happens before  
equilibrium is reached.**



# Swell is governed by propagation



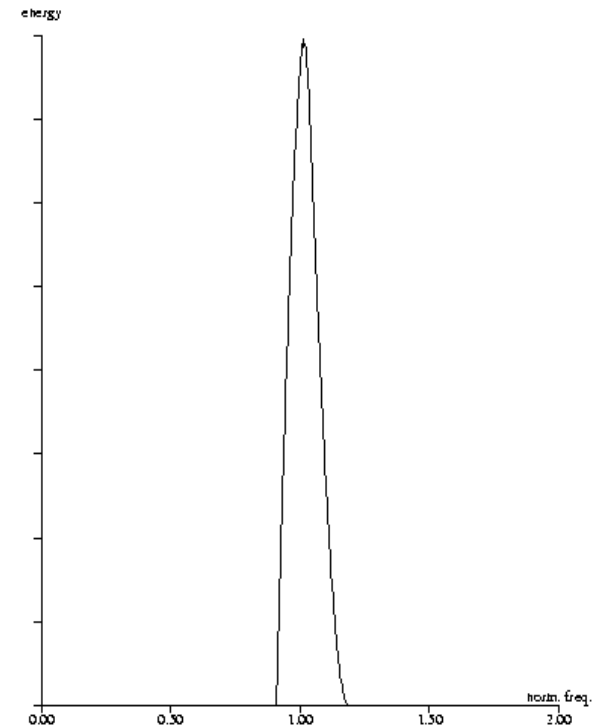
**For swell, we have a  
generating area at some  
time in the past, and  
propagation.**



# Each storm's influence is to be considered separately

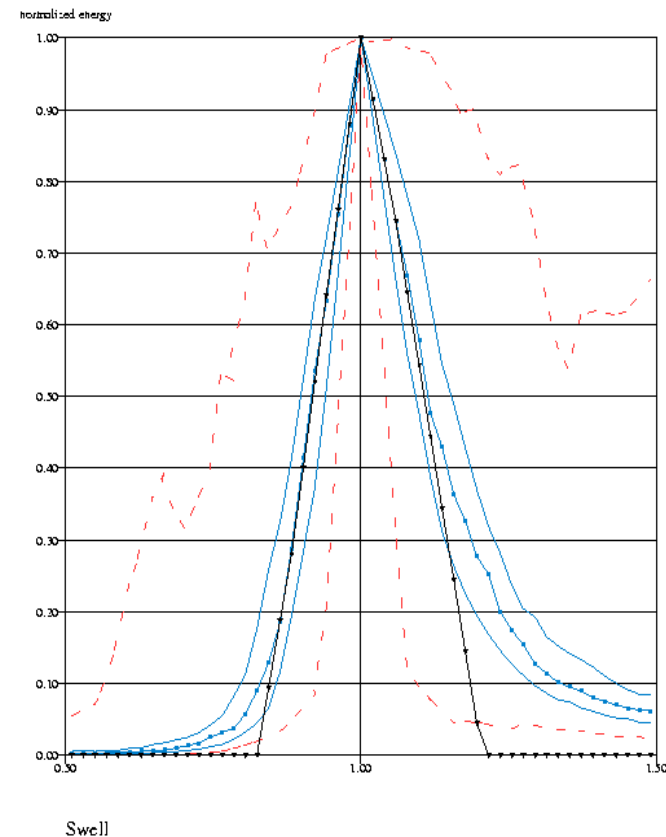
**Propagation carves out  
a shape from the one  
in the generation  
area.**

**Suggestion:  
use a simple spectral  
triangle for each  
system.**



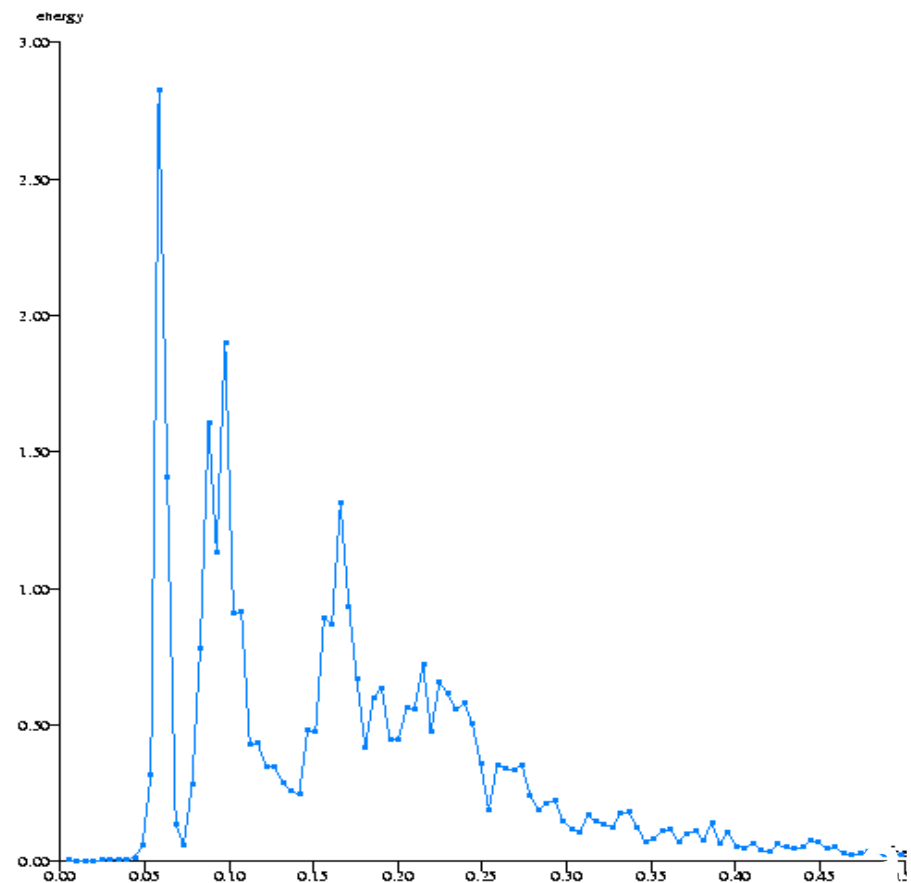
# Triangular shape

- Fitting method from Olagnon (2001).
- Extend from  $(m-1)/m$  fp to  $m/(m-1)$  fp.
- $m$  in the vicinity of 6
- $m$  may need to be increased at some locations.
- Note that  $m$  is related to the peakedness factor (Goda parameter) by  $\varepsilon_{qp} = (4m-2)/3$



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# Extraction



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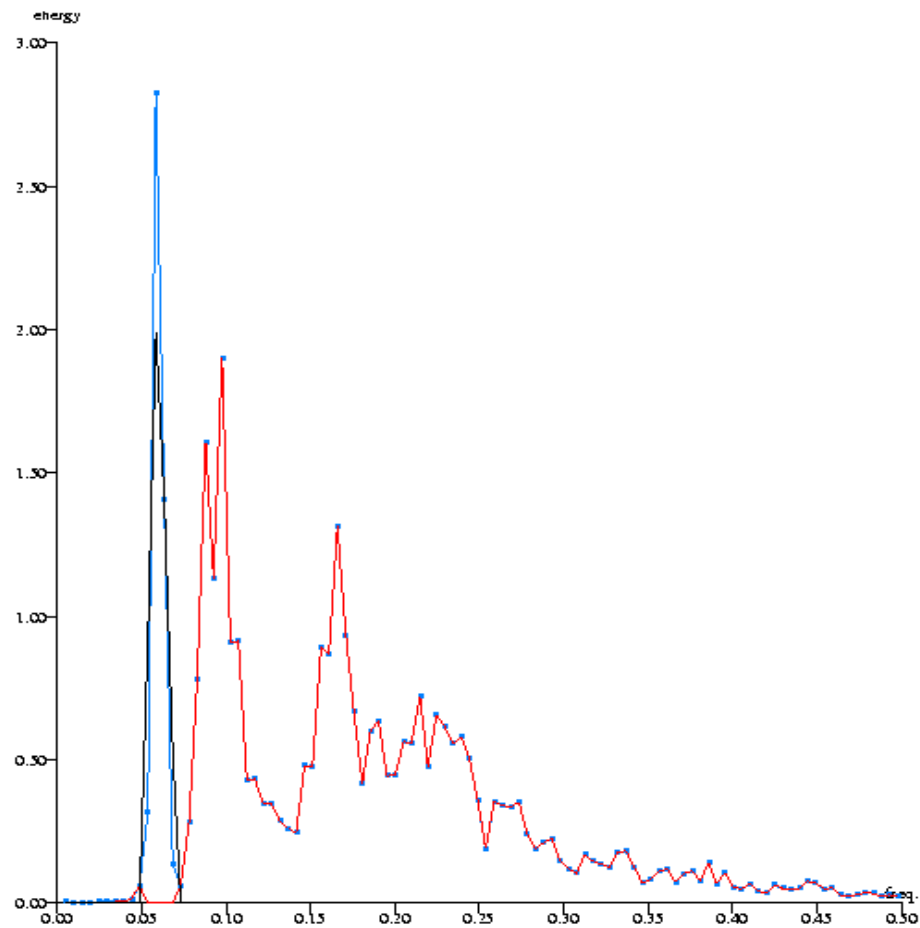
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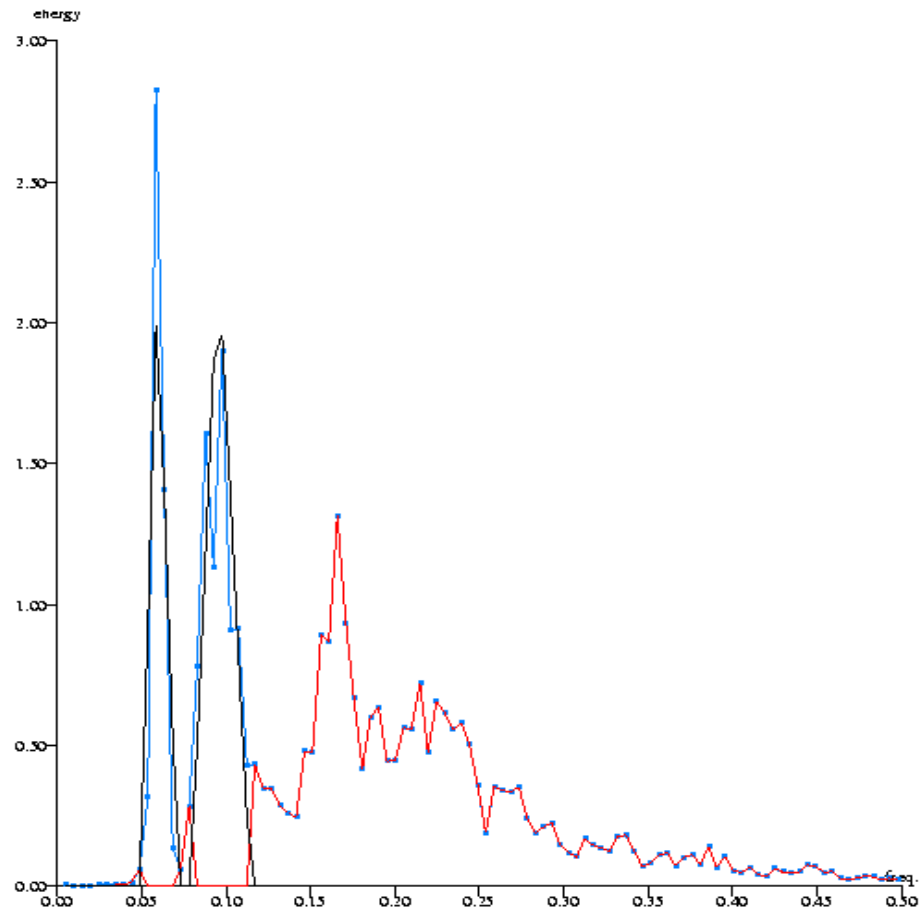


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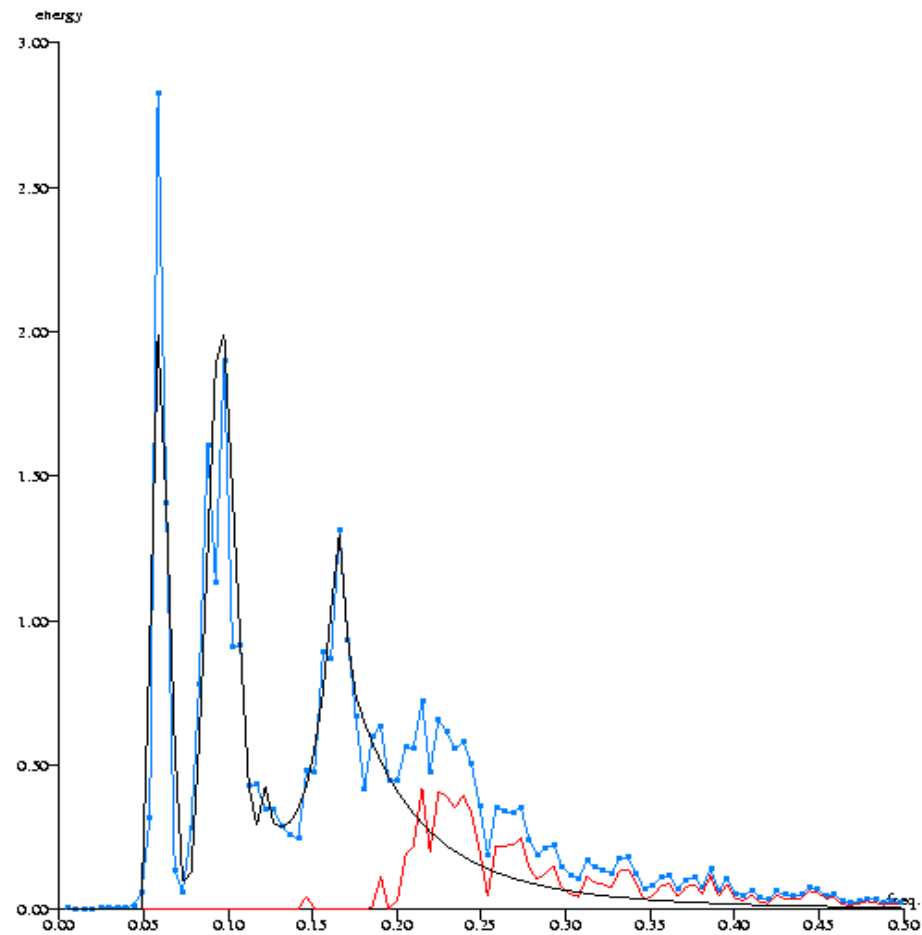


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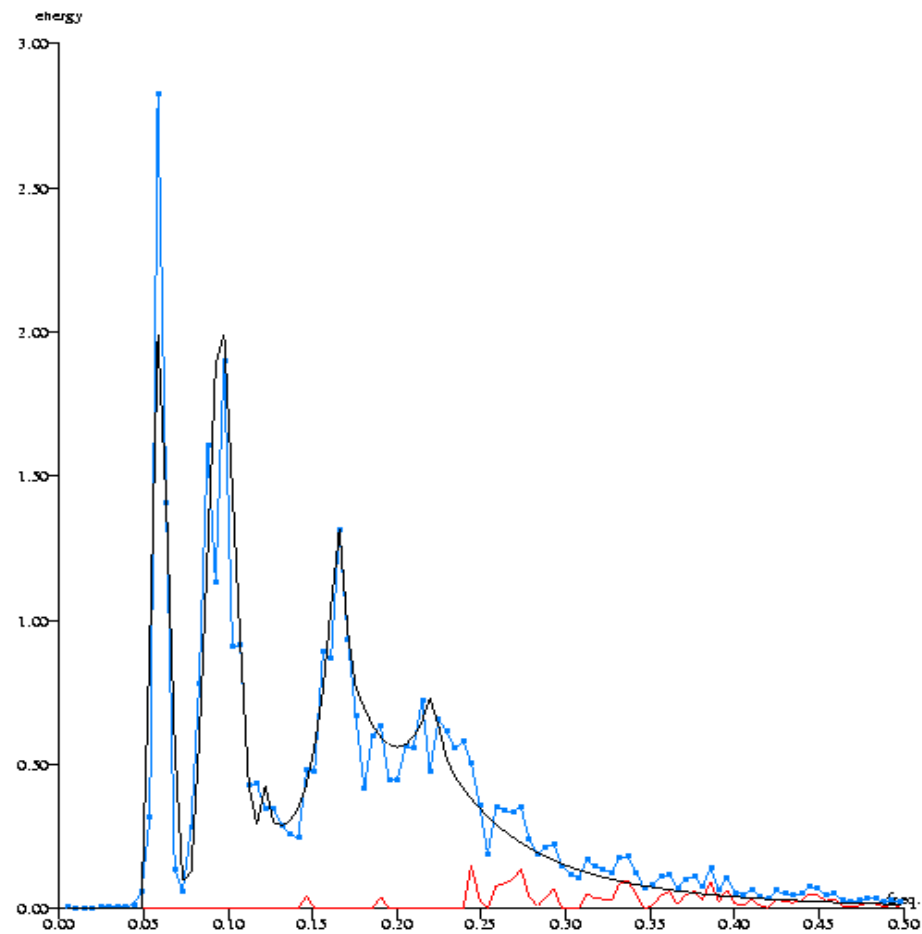


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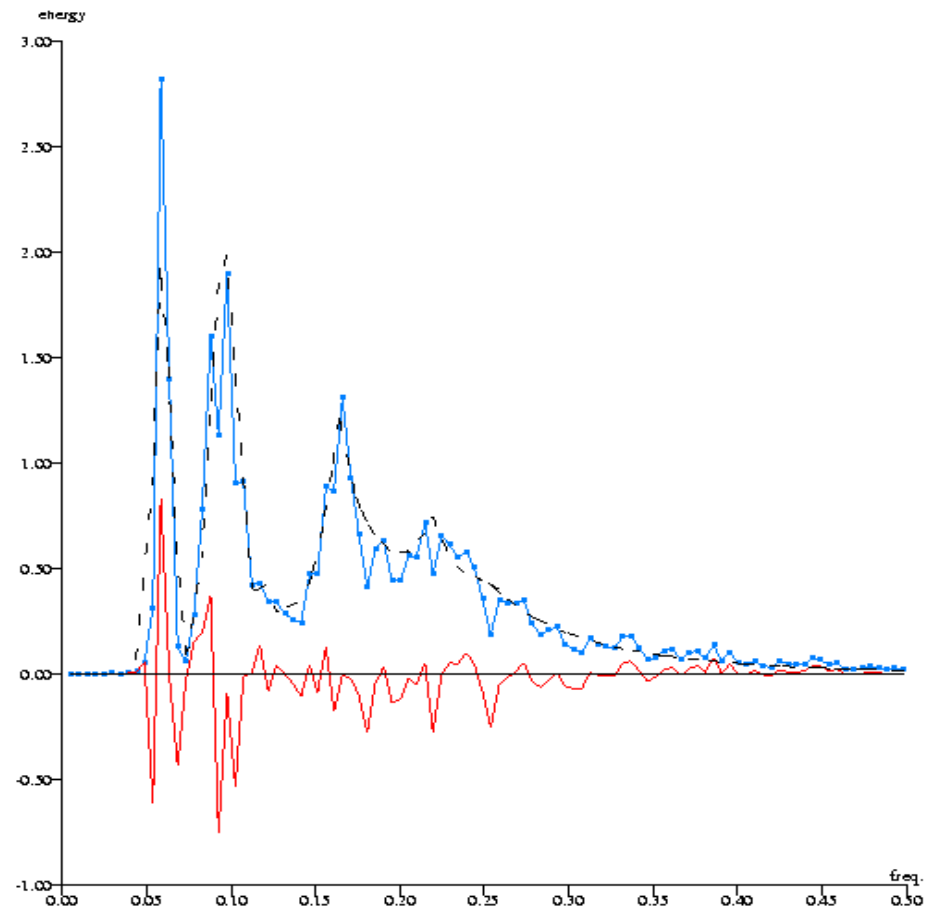


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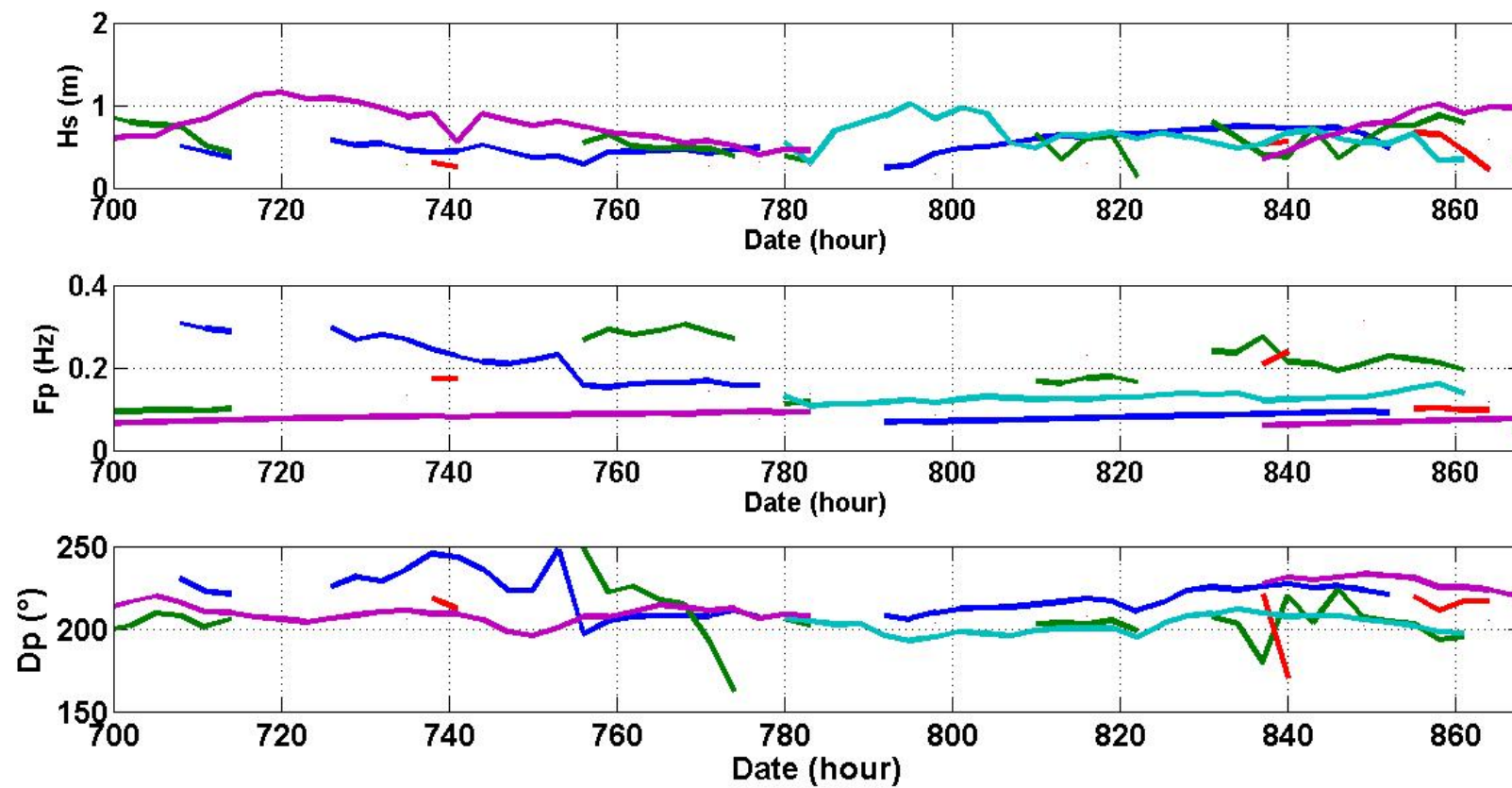
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# Extraction



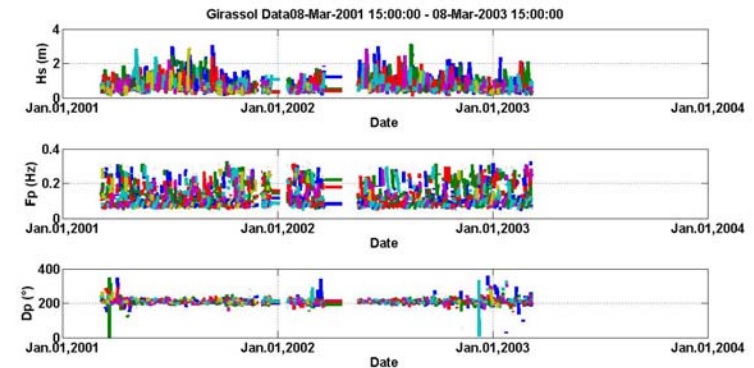


# Systems



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**We have successfully  
replaced a time-history of  
spectra with a time-  
history of a variable  
number of parameters.**



**Now, we can rely on the same “construction” idea and  
method that we used to model spectra from single peaks  
so as to model the process from single wave systems.**

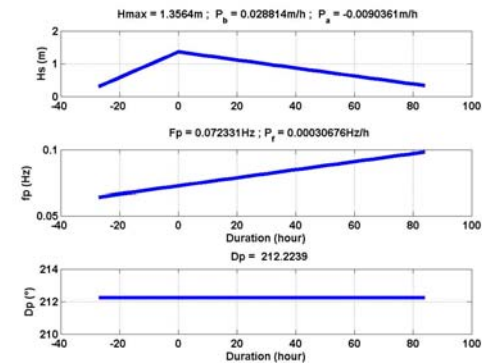
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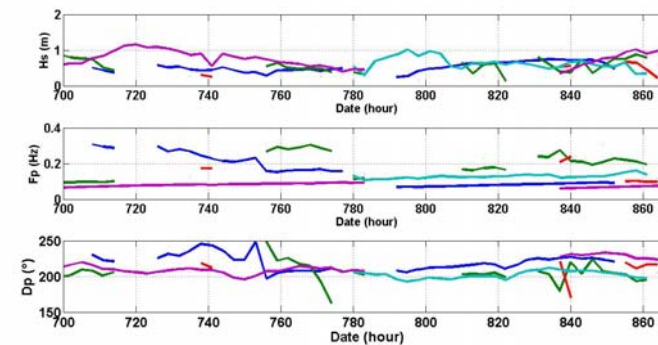
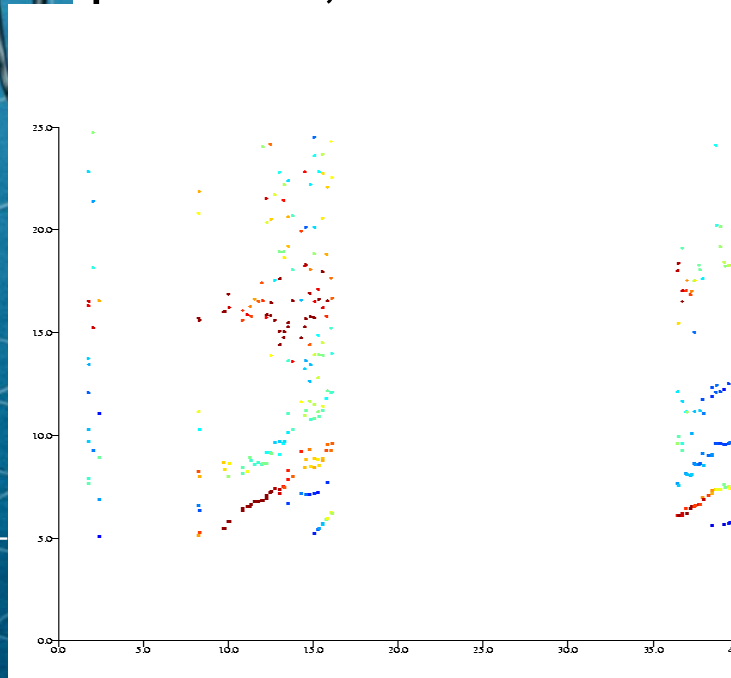
## Let us define an event:

**A climate event is a phenomenon:**

- ✓ **that can be found in all successive observations within a finite, yet significant, duration;**
- ✓ **that can be modeled consistently throughout for each of those observations;**
- **for which the model parameters variations are slow and can be themselves modeled;**
- **and last but not least, that can be traced back to a unique meteorological origin.**



Systems are already “coloured”, i.e. one can follow them over many time-steps, yet some of them may not be pure (at some point, the waves from a new storm are mistaken for the continuation of the swell from an older one), may be short parts of longer events truncated by some measurement or partition problem, etc.

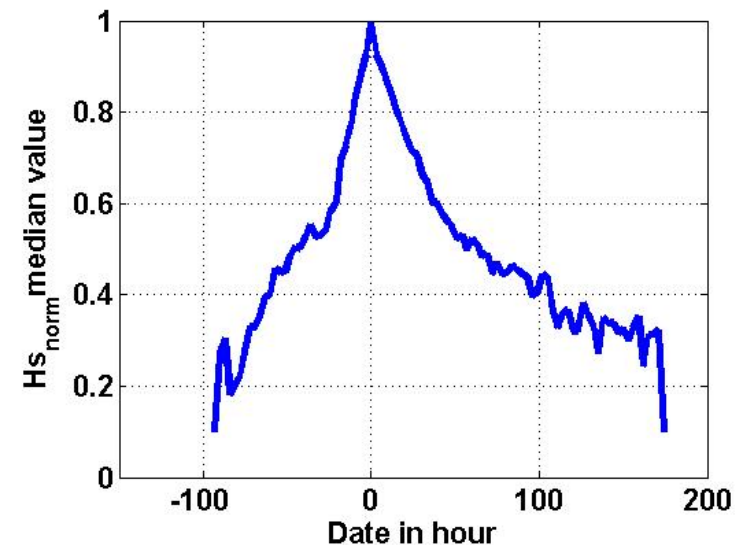
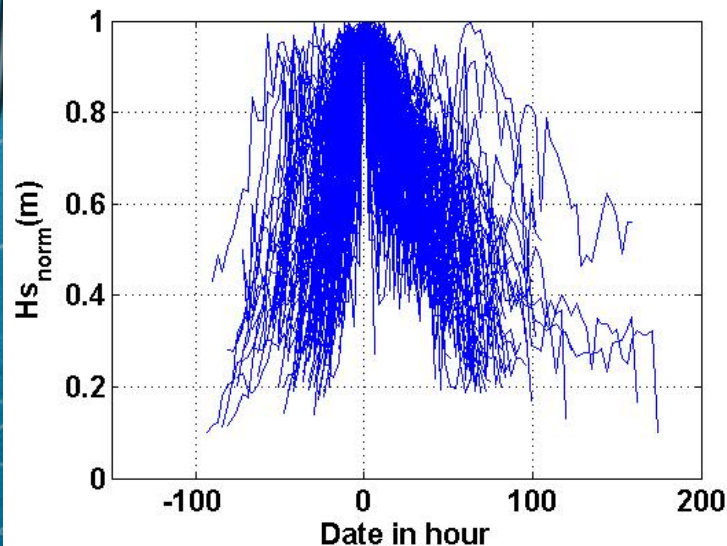


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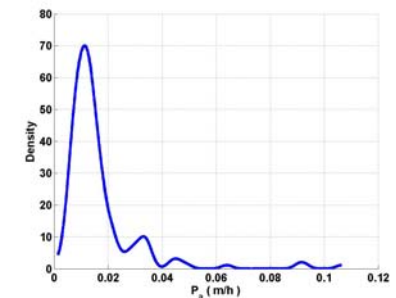
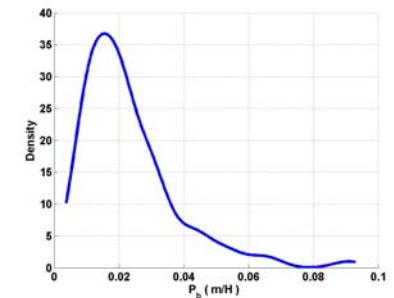
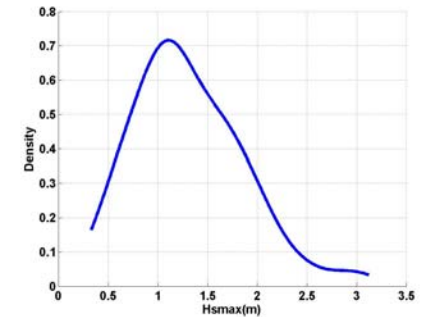
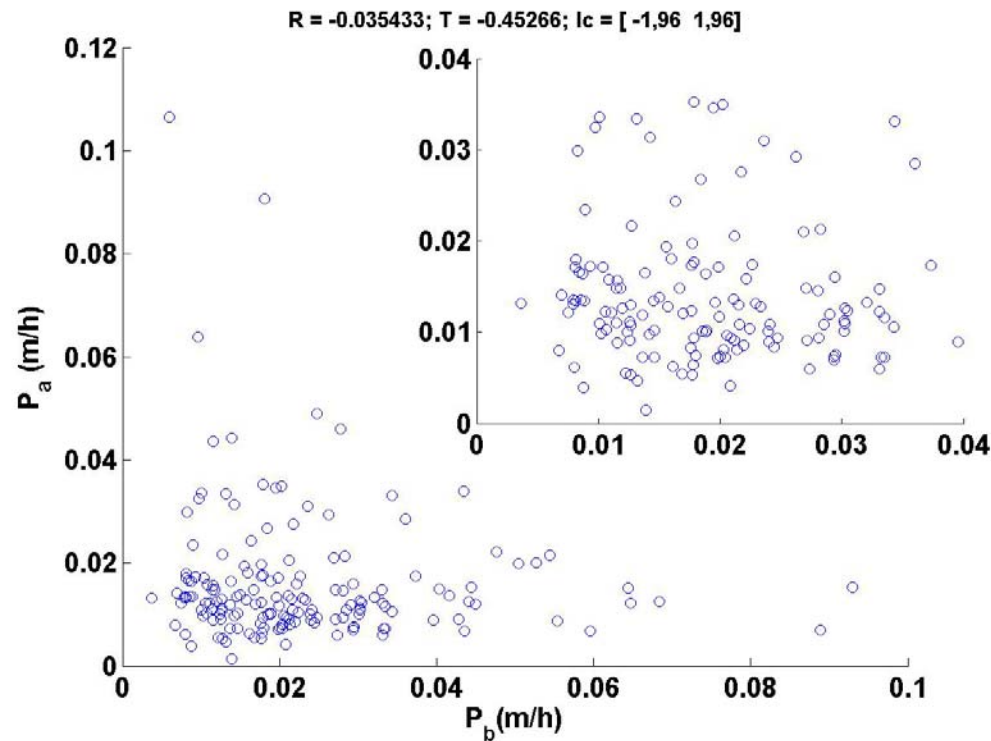


A set of the best events is selected, and a model is sought for their normalized parameters time histories with the same method (Olagnon 2001) as for spectral peaks.



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Hs is thus modeled by Hs\_max of the event, a left  
slope for swell growth, a right slope for swell decay.  
No significant correlations.

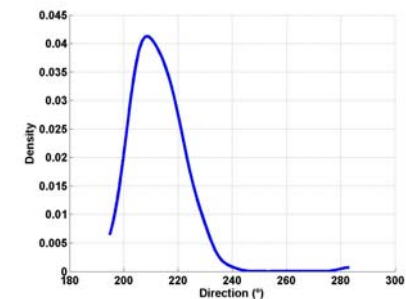
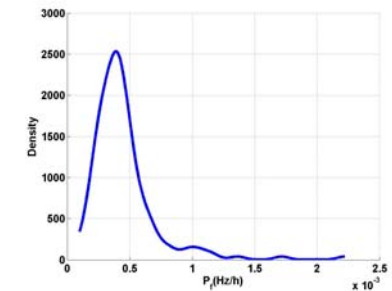
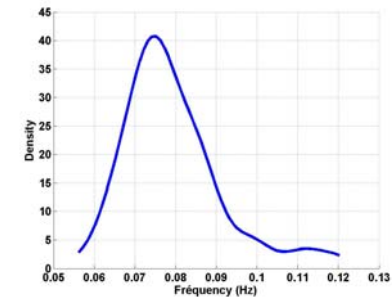
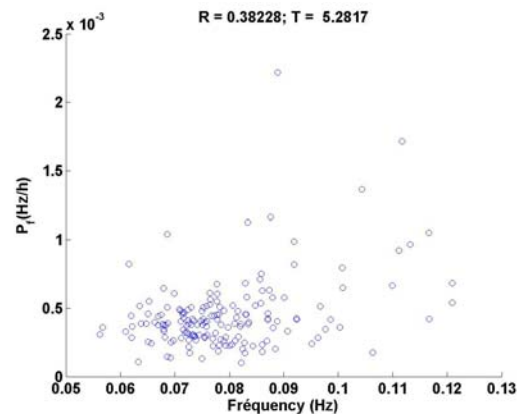
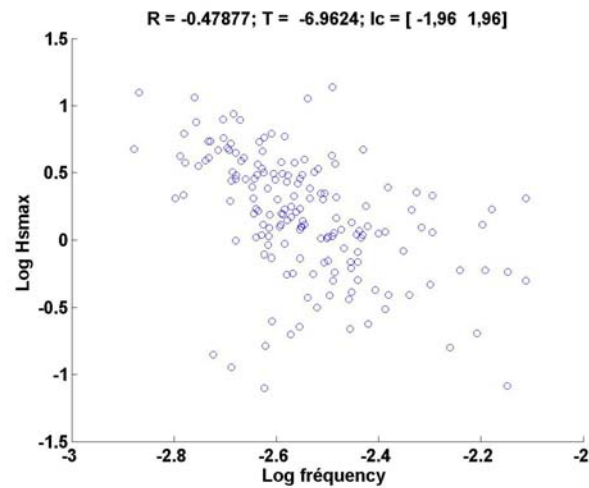


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Frequency is steadily increasing and direction nearly constant, frequency is correlated to  $H_s$  and frequency slope to frequency.

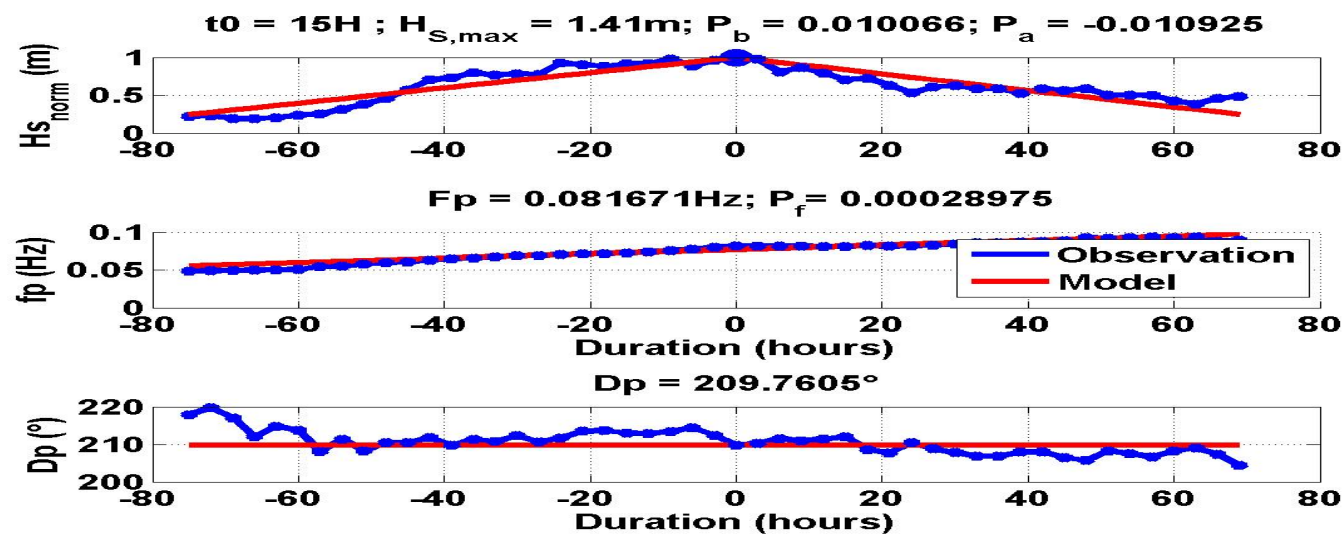


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Thus the following model for an individual swell event:

- $H_s$ : Triangle with growth slope independent of decay slope.
- $F_p$ : Linear increase, with value at  $H_{s\_max}$  dependent on  $H_{s\_max}$ .
- $D_p$ : Constant.



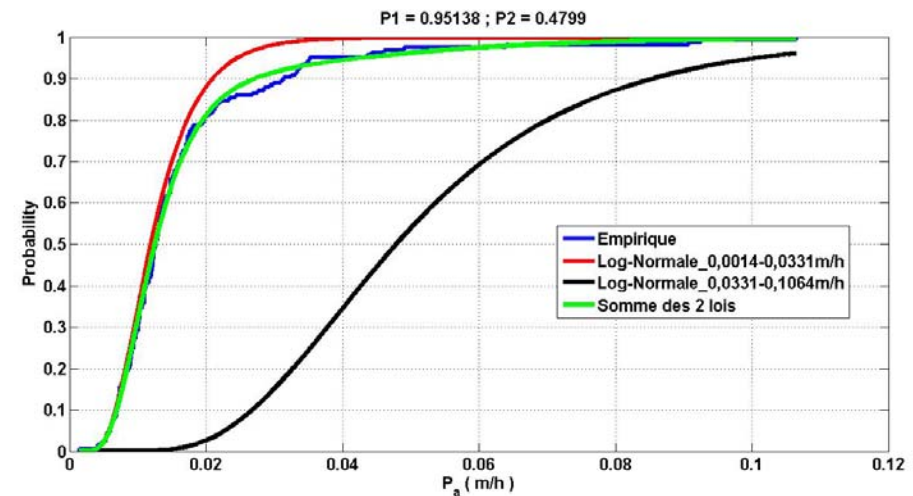
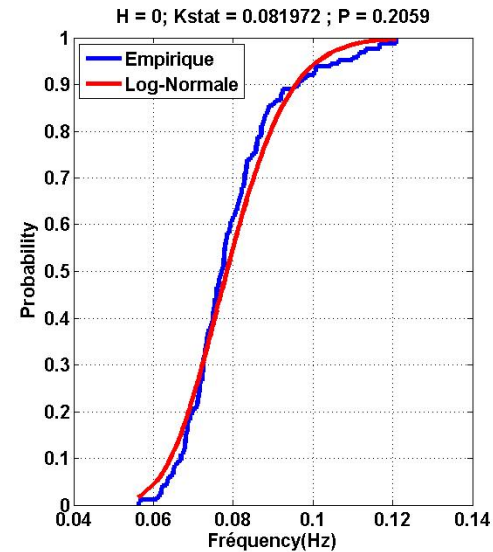
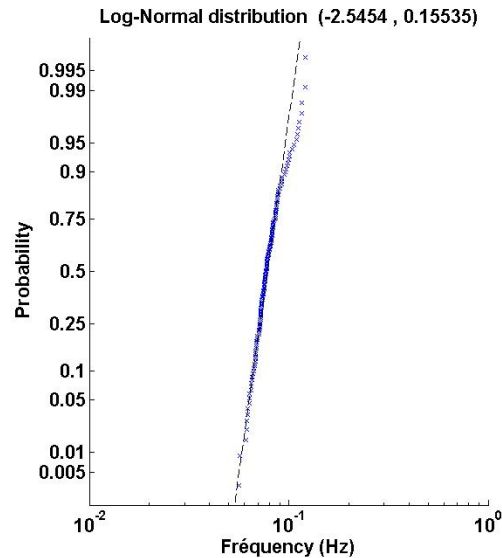


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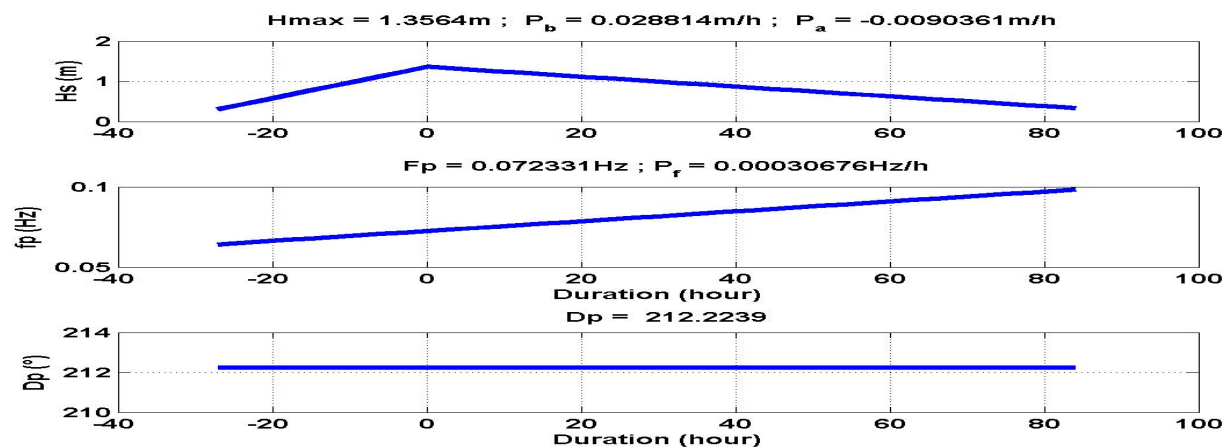
Then we can fit distributions and further investigate correlations for the parameters:  $H_s$ ,  $f_p$ ,  $D_p$ ,  $H_s$  slope left,  $H_s$  slope right,  $f_p$  slope





... and the distributions for the parameters are:

- $H_{s\_max}$ : log-normal distribution.
- Ascending  $H_s$  slope: log-normal distribution.
- Descending  $H_s$  slope: sum of 2 log-normal distributions.
- $F_p$ : log-normal distribution, dependent on  $H_s$ .
- $F_p$  slope: log-normal distribution, dependent on  $F_p$ .
- $D_p$ : 99% truncated normal distribution, with discrete addition. *Most swell systems come from the Southwest sector (South Atlantic), yet on rare but verified instances (about 1%), Northern Hemisphere swells make it to the location where they arrive from the Northwest.*

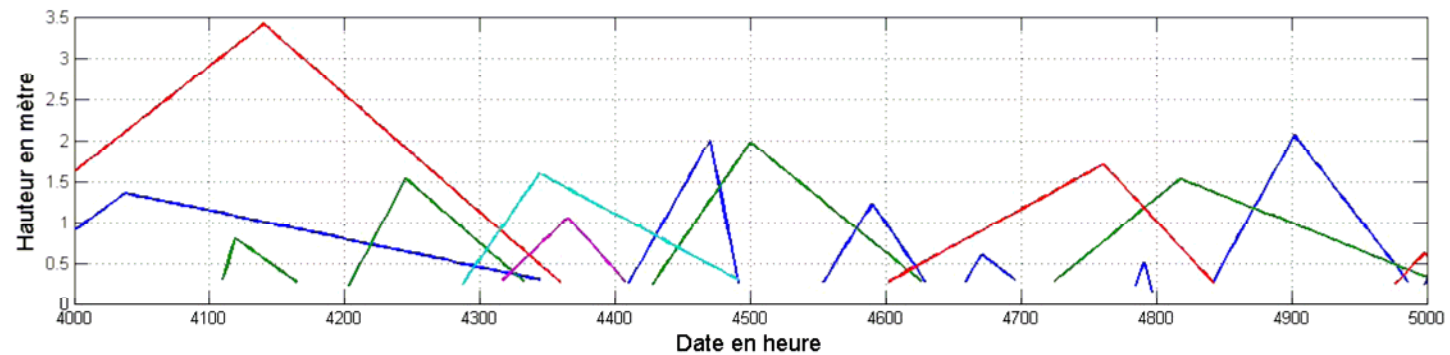
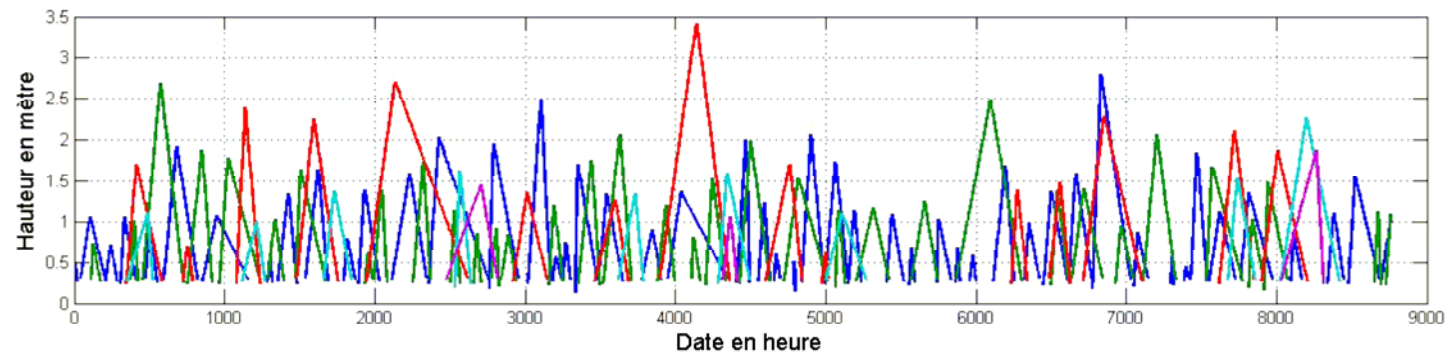


Now we can simulate all the events that would occur within a given duration. We only need to fulfill some condition as to the number of events: since we have only selected “beautiful events”, we don’t know the true occurrence density of events. We impose the condition that the yearly averaged Hs should be the same as the observed one. It has reasonable interannual variability (c.o.v. 8%), so should be correctly estimated over our 2 years of data.

We need a model for the process of the occurrence of events. This is a topic for future research, still we can make a quick and dirty simulation as follows:

- Assume a given distribution shape for the the time-durations between the times of Hs\_max of successive events (for instance, log-normal or sum of 2 log-normals);
- Adjust the parameter(s) of that distribution so as to meet some constraint(s) (for instance, average number of events present at any time = the observed average);
- Draw random independent intervals between events accordingly.

# Reconstructed history



# Example of properties

- **Yearly rms Hs:** assuming times of no swell are measurement failures, database => 1.28m, hindcast on nearby location => interannual c.o.v. 8%
- Reconstruction with target 1.28 conditioned on those sea-states with at least one swell present, inter-event duration weighted sum of 2 log-normals 2.5 and 3.5 days => 1.21m c.o.v. 9%, value 1.28 at fractile 65% of marginal distribution.
- **FPSO Vertical bending moment fatigue damage:** extrapolated to 100 years from the 1.64 validated year of the database => 0.566
- With the above reconstruction of 100 years, damage => 0.529, interannual c.o.v. 55% down to 41% for 1.64 years, value 0.566 at fractile 65% of distribution.



# Conclusion

- If we use a sensible model for the process of swell events rather than the quick and dirty method, we can expect very satisfactory results for almost any application.
- We have developed a method that consists in identifying a model for time-consistent events, and then looking for such events in the data.
- Why not use the same method for the analogs of systems (f.i. eof's) in current profiles ?

